1 Tropical weathering of the Taconic orogeny as a driver for 2 Ordovician cooling

3 **Nicholas L. Swanson-Hysell¹ and Francis A. Macdonald²**

4 *¹Department of Earth and Planetary Science, University of California, Berkeley,*
5 *California 94720-4767, USA*

6 *²Department of Earth and Planetary Sciences, Harvard University, Cambridge,*
7 *Massachusetts 02138, USA*

8 **ABSTRACT**

9   The Earth's climate cooled through the Ordovician Period leading up to the
10 Hirnantian glaciation. Increased weatherability of silicate rocks associated with
11 topography generated on the Appalachian margin during the Taconic orogeny has been
12 proposed as a mechanism for Ordovician cooling. However, paleogeographic
13 reconstructions typically place the Appalachian margin within the arid subtropics, outside
14 of the warm and wet tropics where chemical weathering rates are highest. In this study
15 we reanalyze the paleomagnetic database and conclude that Ordovician constraints from
16 cratonic Laurentia are not robust. Instead, we use paleomagnetic data from well-dated
17 volcanic rocks in the accreting terranes to constrain Laurentia's position given that the
18 Appalachian margin was at, or equatorward of, the paleolatitude of these terranes. To
19 satisfy these allochthonous data, Laurentia must have moved toward the equator during
20 the Ordovician such that the Appalachian margin was within 10° of the equator by 465
21 Ma. This movement into the tropics coincided with the collision and exhumation of the
22 Taconic arc system, recorded by a shift in neodymium isotope data from shale on the

Appalachian margin to more juvenile values. This inflection in detrital neodymium isotope values precedes a major downturn in global seawater strontium isotopic values by more than one million years, as would be predicted from a change in weathering input and the relatively long residence time of strontium in the ocean. These data are consistent with an increase in global weatherability associated with the tropical weathering of mafic and ultramafic lithologies exhumed during the Taconic arc-continent collision. A Taconic related increase in weatherability is a viable mechanism for lowering atmospheric CO₂ levels through silicate weathering contributing to long-term Ordovician cooling.

INTRODUCTION

Ordovician strata record the transition from an Early Ordovician ice-free world to end-Ordovician glaciation and mass extinction (Cooper and Sadler, 2012). Several hypotheses have been proposed to account for this cooling and the initiation of glaciation including: increased carbon burial (Brenchley et al., 1994), aerosol release from volcanism (Buggisch et al., 2010), decreased volcanic outgassing (McKenzie et al., 2014), increased silicate weathering due to topography associated with the Taconic orogeny (Kump et al., 1999), and increased weathering of fresh volcanic rocks (Young et al., 2009). Oxygen isotope data from brachiopods and conodonts indicate that Hirnantian glaciation is the culmination of longer term cooling from 480 to 445 Ma (Trotter et al., 2008; Veizer and Prokoph, 2015). Although short-term perturbations such as increased organic carbon burial inferred from positive carbon isotope excursions, changes in ocean circulation, or sulfur aerosol release could account for transient cooling associated with the Hirnantian glacial maximum, tectonic changes associated with long-term changes to CO₂ sources or sinks are required to drive ~35 m.y. of cooling. An increase in

global weatherability can lead to CO₂ levels decreasing through increased silicate weathering, associated delivery of alkalinity to the ocean, and sequestration of bicarbonate in chemical sediments. The silicate weathering feedback can lead to stabilization at a lower steady-state CO₂ level (Kump et al., 1999).

Arc-continent collision is a tectonic process that can combine the mechanisms for cooling outlined here and lead to a decrease in volcanic outgassing through the death of an arc, and an increase in silicate weathering through increased topography and the exhumation of highly weatherable mafic and ultramafic rocks (Reusch and Maasch, 1998; Jagoutz et al., 2016). Arc-continent collision associated with the Taconic orogeny has been suggested to be associated with Ordovician cooling (Reusch and Maasch, 1998), but paleogeographic reconstructions typically place the Taconic arc system outside of the tropic weathering belt and within the arid subtropics (e.g., Mac Niocaill et al., 1997; Domeier, 2016; Torsvik and Cocks, 2017). Modern evaporite belts and the paleolatitude of evaporites constrain the arid subtropics to be persistently between latitudes of 15° and 35° (Evans, 2006). Given that weathering rates are strongly dependent on temperature and precipitation, and that weathering rates within basaltic watersheds in the tropics are approximately an order of magnitude higher than those in mid-latitudes (Dessert et al., 2003), such a subtropical position would likely preclude major CO₂ drawdown associated with arc-continent collision (Jagoutz et al., 2016). Consequently, the reconstruction of the paleolatitude of the orogeny is critical to the hypothesis that an increase in silicate weatherability associated with the Taconic orogeny drove a portion of Ordovician cooling. Did the Taconic arc-continent collision occur in the arid subtropics or in the wet tropics?

TECTONICS OF THE TACONIC OROGENY

The Taconic orogeny encompasses Ordovician collisional and accretionary events between volcanic arcs that formed within the Iapetus Ocean and the Appalachian margin of Laurentia. The Taconic orogeny has been separated into three broad phases (van Staal and Barr, 2012): Taconic 1 (495–488 Ma) includes local amphibolite-grade metamorphism in the arc terranes; Taconic 2 (488–461 Ma) spans the collision of the leading edge of the Taconic arc system with distended fragments and promontories of the Laurentian margin and the initiation of north-directed subduction (Fig. 1); and Taconic 3 (461–445 Ma) comprises later arc accretion events. By ca. 465 Ma, amalgamated arc terranes and fragments of the margin were thrust onto Laurentia, and delivered arc detritus, including detrital chromite, into marginal basins (e.g., Hiscott, 1978; Macdonald et al., 2017).

The colliding Taconic arc system extended west (paleocoordinates in Fig. 1) into the southern Appalachians as far as Alabama (Hibbard, 2000), and east along the Greenland margin to Ellesmere Island (Trettin, 1987). This elongate east-west exposure of the arc system was all within a similar latitude band (Fig. 1).

PALEOGEOGRAPHY

Concerted efforts over decades of integrating geologic and paleomagnetic data have led to an understanding that from the Cambrian into the Ordovician, Laurentia's Appalachian margin was oriented east-west as the northern boundary of the Iapetus Ocean (Mac Niocaill et al., 1997). Although paleogeographic models typically place this margin south of 20°S in the relatively arid subtropics, this position in the Ordovician is

poorly constrained due to a lack of reliable paleomagnetic poles from cratonic Laurentia. In the comprehensive apparent polar wander path compilation of Torsvik et al. (2012), only two poles are included for the Ordovician: the St. George Group and Table Head Group limestones of Newfoundland. However, the Table Head Group limestones fail a conglomerate test (Hodych, 1989). Therefore, their remanence, and the similar remanence of the underlying St. George Group, must be the result of remagnetization. The Table Head Group rocks pass a fold test, indicating that remagnetization occurred prior to Devonian folding. Exclusion of these poles exacerbates an already large temporal gap between Laurentia poles in the Torsvik et al. (2012) compilation, such that there are no robust poles from the craton between the ca. 490 Ma Oneota Dolomite and the ca. 438 Ma Ringgold Gap poles (Fig. 2). The paleolatitudes implied by Cambrian and Silurian poles for Laurentia's distal margin (e.g., the New York and St. Lawrence promontories) are both in the subtropics (Fig. 2), and extrapolation between these poles (such as a spline fit; Torsvik et al., 2012) keeps Laurentia at a similar position through the Ordovician.

Given that there are no robust Ordovician paleomagnetic data from the Laurentian craton, we take the approach of using paleomagnetic data from well-dated volcanic rocks on the accreting terranes with magnetizations that are interpreted to be primary. Because the Appalachian margin must have been at or equatorward of these terranes, these data provide the best existing constraints on the Ordovician paleolatitude of Laurentia and have been interpreted to indicate the presence of peri-Laurentian, intra-Iapetan, and peri-Avalonian arc volcanism (Mac Niocaill et al., 1997). Open source reconstructions developed in GPlates software (<https://www.gplates.org/>) for the

115 evolution of the Iapetus Ocean (Domeier, 2016; Torsvik and Cocks, 2017) provide an
116 excellent framework that can be modified with this approach.

117   In contrast to the Laurentian craton, eight robust Ordovician paleomagnetic data
118 sets have been reported from accreted Taconic arc terranes through extensive efforts of
119 the Rob Van der Voo research group at the University of
120
121 Michigan (USA) (see the GSA Data Repository¹). The interpretation of primary
122 remanence in these volcanic rocks is variably based on dual polarities, positive fold tests,
123 and interpretation of magnetic mineralogy. The oldest such locality is within the Notre
124 Dame arc of Newfoundland, where ca. 477 Ma mafic volcanics of the Moreton's Harbour
125 Group yielded a paleolatitude of $\sim 11^{\circ}\text{S}$
126 (8° – 15°S at 95% confidence) and were therefore interpreted to have formed in close
127 proximity to Laurentia (Johnson et al., 1991). Four paleomagnetic localities from ca.
128 470–465 Ma volcanic rocks of the Victoria arc of Newfoundland provide paleolatitude
129 constraints; the lowest latitude results are from the Lawrence Head volcanics,
130 which were at $\sim 12^{\circ}\text{S}$ (2° – 24°S at 95% confidence) (see
131 the Data Repository). Similar aged pillow lavas from arc terranes in Newfoundland (the
132 Annieopsquotch arcs in Fig. 2) give paleolatitudes of $\sim 30^{\circ}\text{S}$ that have been interpreted to
133 indicate that they formed some distance from the margin within the Iapetus Ocean (Van
134 der Voo et al., 1991). In New England (northeastern United States), ca. 467 Ma volcanics
135 of the Bronson Hill arc yield a paleolatitude of $\sim 20^{\circ}\text{S}$ (12° – 29°S at 95% confidence)
136 while younger ca. 458 Ma volcanics give paleolatitudes of $\sim 14^{\circ}\text{S}$ (8° – 23°S at 95%
137 confidence) and $\sim 11^{\circ}\text{S}$ (6° – 16°S at 95% confidence).

Although the Notre Dame arc was at a low latitude by ca. 475 Ma (Johnson et al., 1991) when it collided with hyperextended fragments of the Laurentian margin (Macdonald et al., 2014; van Staal and Barr, 2012), the Taconic seaway separated these terranes from the Laurentian autochthon until they were exhumed ca. 465 Ma. While the width of the Taconic seaway is unconstrained, the hyperextended margin of northeast Australia, which extends >500 km from the craton, may be a modern analog. This ~500-km-wide seaway closed between 475 and 465 Ma.

Paleolatitude constraints from ca. 470–465 Ma volcanics of the Taconic arc terranes span ~20° of latitude, suggesting a distended arc system comparable to the modern southwest Pacific arc system (Fig. 1; Mac Niocaill et al., 1997). Although the precise latitudinal spread is difficult to resolve given uncertainty associated with paleolatitude estimates, we interpret the spread of these latitudes to represent the leading and trailing edges of the arc system (Fig. 1). This approach is a simplification; analogous to the modern southwest Pacific, there were probably other active subduction zones. Shortening during the Taconic and subsequent orogenies would have translated these terranes inward toward the craton, further contributing to the interpretation that Laurentia was equatorward of their paleolatitudes. Overall, the paleomagnetic database strongly supports a revised reconstruction wherein the Appalachian Laurentian margin was equatorward of 10°S at 465 Ma (Fig. 1).

WEATHERING PROXY DATA

Strontium and neodymium isotope data were compiled and recalculated (see the Data Repository) using *The Geological Time Scale 2012* (see Cooper and Sadler, 2012).

⁸⁷Sr/⁸⁶Sr data developed from the conodont apatite record a broad decline from 0.7090 to 0.7088 between 480 and 465 Ma. This gradual decline is followed by a sharp deflection at 465 Ma toward more juvenile ⁸⁷Sr/⁸⁶Sr values, reaching 0.7079 by 450 Ma (Saltzman et al., 2014; Fig. 2). Neodymium isotope (ϵ_{Nd}) data from fine-grained siliciclastic rocks deposited on the distal margin of Laurentia, on the Taconic allochthon, and Sevier basin (Gleason et al., 2002; Macdonald et al., 2017) display an inflection to more positive values at 465 Ma consistent with a substantial increase of sediment being weathered from juvenile lithologies (Fig. 2). This inflection in ϵ_{Nd} values occurs later in more interior basins (Fig. 2) that did not receive arc-derived sediment until subsequent accretionary events thrust arc rocks onto Laurentia between ca. 455 and 450 Ma (Macdonald et al., 2014, 2017).

DISCUSSION

The paleogeographic reconstruction presented here suggests that the Appalachian margin was at a significantly lower latitude than is typically depicted, equatorward of 10°S by 465 Ma (Fig. 1). Our reconstruction is compatible with paleomagnetic data from the Taconic arc system and is not in conflict with robust paleomagnetic poles from Laurentia.

We propose that the broad rise in oxygen isotope values and decline in strontium isotope values between 490 and 465 Ma (Fig. 2) are related to the movement of the Taconic arc system into the tropics and collision of the leading edge with distended fragments and promontories of the Laurentian margin (Taconic orogenic phase 2 of van Staal and Barr, 2012). A concomitant increase in global weatherability would have caused cooling through CO₂ drawdown, moderated by the silicate weathering feedback.

184 In addition, we argue that the sharp drop in $^{87}\text{Sr}/^{86}\text{Sr}$ values, the shift toward more
185 juvenile ϵ_{Nd} values in shale from the distal margin of Laurentia, and the additional
186 increase in oxygen isotope values between 465 and 455 Ma (Fig. 2) are due to the uplift
187 and exhumation of the Taconic arc system in the tropics (peak of Taconic 2) followed by
188 continued Late Ordovician arc accretion (Taconic 3). This exhumation led to uplift and
189 erosion of island arc volcanics and suprasubduction ophiolites, as evidenced by the
190 presence of detrital chromite in Middle to Late Ordovician foreland basins (e.g., Hiscott,
191 1978).

192 Increased weathering of volcanic arcs associated with the Taconic orogeny was
193 previously invoked to explain the Ordovician drop in $^{87}\text{Sr}/^{86}\text{Sr}$ values (Young et al.,
194 2009). The feasibility of this scenario was supported with a model in which global
195 weatherability was increased by 25% and a new flux of riverine $^{87}\text{Sr}/^{86}\text{Sr}$ was introduced
196 from weathering basalt with a composition of 0.7043 (Young et al., 2009). The ϵ_{Nd}
197 compilation from the Appalachian margin of Laurentia, which records local provenance,
198 is consistent with the hypothesis that the Taconic orogeny played a significant role in the
199 inferred increase in global weatherability and riverine $^{87}\text{Sr}/^{86}\text{Sr}$ input to the ocean. The
200 inflection in ϵ_{Nd} data from distal margin basins occurs a few million years prior to the
201 inflection in the global $^{87}\text{Sr}/^{86}\text{Sr}$ curve (Fig. 2). This lead time is predicted if the
202 weathering of Taconic terranes is a significant driver of the global strontium signal.
203 Juvenile ϵ_{Nd} values should be imparted in siliciclastic rocks over the time scale that
204 sediment transits from source to sink (thousand year time scales), whereas strontium has
205 a multimillion year residence time in the ocean such that a prolonged interval of arc
206 weathering would be necessary to significantly change seawater $^{87}\text{Sr}/^{86}\text{Sr}$. A complication

in this interpretation is that the age model for the ϵ_{Nd} data is anchored by U-Pb zircon ages from ashes within the same stratigraphic sections (Macdonald et al., 2017), whereas the $^{87}\text{Sr}/^{86}\text{Sr}$ age model is based on Cooper and Sadler (2012; Saltzman et al., 2014), so the estimated temporal offset is as accurate as the calibration of the geological time scale.

Although other arc systems likely enhanced global weatherability in the Ordovician, such as those in the paleo-Asian Ocean and the Fammetanian arc of present-day Argentina, the Taconic arcs likely played an outsized role as they were exhumed along an east-west belt in the tropics during the closure of the Iapetus Ocean (Fig. 1). Exhumation would have created significant topography composed of mafic and ultramafic lithologies through a wide swath across the tropics. This scenario has similarities to the low-latitude closure of the Neo-Tethys Ocean, and two-phase collision of the trans-Tethyan subduction system, which coincided with the two-pronged cooling trend from the Cretaceous to Oligocene (Jagoutz et al., 2016). The closure of major oceanic basins along east-west belts in the tropics may have been a significant driver of long-term cooling trends throughout Earth history. Following the Taconic orogeny, the Appalachian margin moved away from the tropics, so that collisions associated with the Salinic orogeny in the Silurian would have occurred at $\sim 20^\circ\text{S}$, where there would have been a lesser effect on global weatherability (Fig. 2).

Lower $p\text{CO}_2$ resulting from elevated global weatherability could have set the stage for the growth of ice sheets during the Hirnantian. However, these tectonic boundary conditions may not be the sole driver for the Hirnantian ice advance, and other factors such as orbital forcing, changing ocean circulation, organic carbon burial, or rapid

changes in albedo may have caused the shorter term cooling associated with the Hirnantian glacial maximum.

CONCLUSIONS

Our paleogeographic reconstruction constrained by the paleolatitude of allochthonous volcanic rocks demonstrates that Laurentia moved toward the equator during the Ordovician such that the Appalachian margin was equatorward of 10°S at 465 Ma. This movement into the tropics coincided with (1) collision and exhumation of the Taconic arc system marked by the appearance of detrital chromite in foreland basins; (2) a shift in ϵ_{Nd} data from fine-grained siliciclastic rocks on the Laurentian margin to more juvenile values; (3) a drop in seawater $^{87}\text{Sr}/^{86}\text{Sr}$ values to more juvenile values; and (4) a continued trend to higher values in the oxygen isotopic composition of both brachiopod carbonate and conodont phosphate. These data are consistent with tropical weathering of the Taconic arc-continent collision as a driver of Ordovician cooling.

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345 **FIGURE CAPTIONS**

346 Figure 1. Paleogeographic reconstruction ca. 465 Ma, after the arrival of the leading edge
347 of the Taconic arc system in the tropics along with the paleolatitude from allochthonous
348 volcanic rocks shown with 95% uncertainty. The reconstructed positions of these
349 paleomagnetic localities are shown on the classic position of Laurentia (as in Torsvik and
350 Cocks, 2017) and the new position proposed herein. While Laurentia most have been
351 north of these volcanics, in the classic reconstruction their positions
352 are south of the paleolatitudinal constraints
353 rather than equatorward, as in the revised position. The positions of other continental
354 blocks are as in Torsvik and Cocks (2017), other than Carolina, which is modified to be
355 traveling in unison with Ganderia.

356

357 Figure 2. Paleomagnetic and geochemical data from 500 to 400 Ma. A: Paleolatitude
358 constraints for Laurentia, Taconic arc terranes (Popelogan-Victoria, Bronson Hill,
359 Annieopsquotch, and Notre Dame), and the peri-Gondwana Ganderia and Avalonia
360 terranes. Laurentia paleolatitudes are calculated for two localities on the margin from
361 paleomagnetic poles with the implied position of New York (NY) shown for the classic

and new models. B: Strontium isotope data from conodont apatite and brachiopod calcite with a locally weighted scatterplot smoothing (LOWESS) regression curve to the data of Saltzman et al. (2014). C: Neodymium isotope data from fine-grained siliciclastic rocks on the Appalachian margin of Laurentia with a LOWESS curve for distal margin data. D: Oxygen isotope data from conodont apatite and brachiopod calcite with a LOWESS curve for the brachiopod data. VPDB—Vienna Peedee belemnite; VSMOW—Vienna standard mean ocean water. E: Orogenic phases wherein Taconic 2 spans the collision of the leading edge of the arc system with promontories of the Laurentian margin. The peak of Taconic 2 coincides with arc exhumation in the tropics and weathering of ophiolite and arc detritus into Laurentian foreland basins. Late Ordovician arc accretion composes Taconic 3. Data sources are provided in the Data Repository (see footnote 1).

¹GSA Data Repository item 2017238, details of the paleomagnetic and chemostratigraphic data compilations, is available online at <http://www.geosociety.org/datarepository/2017/> or on request from editing@geosociety.org.

Notre Dame arc	Moreton's Harbour Volcanics (477.4 ± 0.4 Ma)
	- paleolatitude - position in classic reconstruction - position in new reconstruction
Popelogan-Victoria arc	Lawrence Head Volcanics (465 ± 2 Ma)
	- paleolatitude - position in classic reconstruction - position in new reconstruction
Bronson Hill arc	Winterville Volcanics (456.5 ± 3.5 Ma)
	- paleolatitude - position in classic reconstruction - position in new reconstruction
	Bluffer Pond Volcanics (456.5 ± 3.5 Ma)
	paleolatitude
	Stacyville Volcanics (467 ± 5 Ma)
	paleolatitude
	other arc volcanics (<i>Annieopsquotch</i> and <i>Popelogan-Victoria</i> ; ca. 468 Ma)
	paleolatitude